

Nitrous oxide emissions from a New Zealand cropped soil: tillage effects, spatial and seasonal variability

M.A. Choudhary^a, A. Akramkhanov^{b,1}, S. Saggar^{c,*}

^a Institute of Technology and Engineering, Massey University, Palmerston North, New Zealand

^b Tashkent Institute of Irrigation and Agricultural Mechanisation Engineers, Tashkent, Uzbekistan

^c Landcare Research, Private Bag 11052, Palmerston North, New Zealand

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Abstract

Agricultural practices are believed to be the major anthropogenic source of enhanced nitrous oxide (N₂O) gas emissions in New Zealand. Studies conducted in New Zealand generally suggest low N₂O emission from pasture; however, there is little information for arable farming systems. This paper evaluates tillage and land use effects on N₂O emissions using a closed chamber technique at an Ohakea silt loam (Gleyic luvisol) where winter oats (*Avena sativa* L.)/fodder maize (*Zea mays* L.) was double-cropped for 5 years. The tillage types included conventional tillage (CT) and no-tillage (NT) systems, and a permanent pasture (PP) was used as a control.

Spatial variability in all treatments showed large inherent variations in N₂O fluxes (a mean CV = 119%), which reflected natural soil heterogeneity, and perhaps the measurement technique used rather than the real differences due to the tillage and cropping systems evaluated. On an annualised basis, N₂O emissions measured from December 1998 to September 1999 from the PP (1.66 kg N₂O-N/ha per year or 19 µg N₂O-N/(m² h)) were significantly lower than the CT and NT fields averaging at 9.20 (or 105) and 12.0 (or 137) kg N₂O-N/ha per year (or µg N₂O-N/(m² h)), respectively. However, there were no differences in N₂O emission rates between the CT and NT treatments. Seedbed preparation using a power harrow which followed within a few days of first ploughing the CT field reduced N₂O emissions by 65% within the first hour after power harrowing. However, N₂O emission rates returned to the pre-power harrowing levels at the next sampling period, which was 1 month later.

There was a strong relationship between log-transformed data of soil water content (SWC) and N₂O emissions in all treatments with $r = 0.73, 0.75$ and 0.86 for the PP, CT and NT treatments, respectively. Seasonal variations in N₂O emission from the PP were in the order of winter = autumn > summer. Although fluxes in the CT were higher in winter than in the autumn season, there were no differences between the summer and autumn data. The seasonal variations in N₂O emission in the NT treatment were in the order of winter > autumn = summer.

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1. Introduction

There is a growing concern worldwide about climate change. Atmospheric warming which is known to be caused by “greenhouse gases” mainly include carbon dioxide (CO₂), methane (CH₄) and

* Corresponding author. Tel: +64-6-356-7154;
fax: +64-6-355-9230.

E-mail address: saggar@landcare.cri.nz (S. Saggar).

¹ Formerly at Institute of Technology and Engineering, Massey University, Palmerston North, New Zealand.

nitrous oxide (N₂O) and to a lesser extent chlorofluorocarbons (CFCs) (IAEA, 1992). The increase in greenhouse gases other than CO₂ in changing the climate is similar in importance to CO₂. Nitrous oxide, which despite its low concentration in the atmosphere of about 310 ppb (IPCC, 1995), on a molecular basis has a radiative force of about 200–300 times that of CO₂ (Jaques, 1992) and an average atmospheric lifetime of about 150 years (IAEA, 1992).

It is widely accepted that the main anthropogenic source of N₂O is agriculture and most N₂O originates as an intermediate product from microbial nitrification and denitrification (Delwiche, 1981). Increased N₂O emissions from soils are associated with fertilisation of soils with mineral nitrogen (N), animal manure, N derived from biological N₂ fixation, and enhanced N mineralisation (MacKenzie et al., 1998). With fertiliser usage predicted to grow worldwide at 6–7% per annum (Peoples et al., 1995) and low N utilisation efficiency in agricultural systems, the potential of soils to form and emit N₂O is expected to markedly increase. Such increases in N₂O emission to the atmosphere are of great concern and need quantification.

Amounts of N₂O emitted depend on complex interactions between soil properties, climatic factors and agricultural practices (Granli and Bockman, 1994). Main factors in the soil controlling N₂O emissions are soil content of NH₄⁺ and NO₃⁻ (Ball et al., 1997; Castaldi and Smith, 1998; Seneviratne and van Holm, 1998); soil aeration status and soil water content (SWC) (Carran et al., 1995; Teira-Esmatges et al., 1998; MacKenzie et al., 1998); presence of degradable organic material which promotes microbial activity (Ineson et al., 1998; Kaiser et al., 1998); soil pH and soil temperature (Mahmood et al., 1998). Although these are known interacting factors they are not always strongly correlated with N₂O fluxes. Due to complexity of interactions between various factors, N₂O emissions have very high spatial and temporal variations.

Intensive use of cultivation practices hugely impacts soil properties, which may influence the extent of N₂O emissions. In the Manawatu region continuous cropping of heavier textured soils results in loss of soil organic matter (SOM) and deterioration in soil structure (Shepherd et al., 2001). Even short-term tillage operations can affect SOM levels and microbial

biomass (Aslam et al., 1999), which are important in nutrient transformations. Conversion of pastures to arable cropping in New Zealand results in depletion of SOM and soil N fertility over time (Saggar et al., 2001) and additional N fertilisers are applied to compensate for the loss of organic N reserves. Since N fertilisation as a source of N is considered to be a major source of increasing N₂O emission, careless N application on such soils may contribute to increased loss of N either as NO₃⁻ through leaching or as N₂O emissions.

An alternative to conventional tillage (CT) is no-tillage (NT) that aims for sustainable agricultural production. A growing acceptance of NT is largely due to reduced soil erosion and runoff (Choudhary and Baker, 1993), enhanced water retention and infiltration (Choudhary et al., 1996), lower summer temperatures (Prihar et al., 1996) and possible increased net return to the farmer (Reicosky, 1994). However, the impact NT has on N₂O emission is not known for these Manawatu soils.

In the past 20 years, research on N₂O emissions has concentrated on enhancing our understanding of N₂O production processes and its controlling factors. Despite this, it is not possible to predict the fate of a unit of N that is applied on a specific arable field (Mosier et al., 1996). Both short- and long-term in situ measurements are needed to assess N₂O emissions from soils. Studies by Ruz-Jerez et al. (1994) and Carran et al. (1995) from both poorly and well drained grazed pastures in the Manawatu region suggest low N₂O emission from these low fertility hill lands. However, there is little information for arable farming systems. Therefore, there is a need for site-specific assessment of the impact of tillage practices on N₂O fluxes.

To evaluate impacts of different tillage systems and land use on N₂O emissions, selected farming practices were chosen in the Manawatu region of New Zealand. The overall aim was to characterise land use practices and their effect on N₂O emissions. The specific objectives of this study were: (a) to quantitatively determine the rates of N₂O emissions from the CT and NT cropping regimes and compare these with permanent pasture (PP) fields throughout one management cycle, and (b) to measure the response of soil N₂O emissions to changes in soil water and soil temperature.

2. Materials and methods

2.1. Experimental site

The experimental site was at the Massey University, Turitea campus (latitude 40°21'S, 175°39'E), on an Ohakea silt loam soil (Typic andoaqualf) classified as Gleyic luvisol (FAO). The treatments included CT, NT and rotationally sheep grazed PP used as a control (established since 1995). The CT included one mouldboard ploughing (to about 25 cm depth) followed by rolling with a heavy roller, and two power harrowings (to about 15 cm depth) at suitable intervals to prepare the seedbed. The NT used spraying with 4 l of glyphosate per ha and drilling with a NT seed drill. Each treatment had four replicates in a randomised block design. Each plot was 17 m long and 3.6 m wide with a 5 m headland for machinery operation on both sides of the plots. In November 1998, a summer crop of fodder maize (*Zea mays* L.) at seed rate of 65 kg/ha and in April 1999 a winter crop of oats (*Avena sativa* L.) at seed rate of 120 kg/ha were sown on CT and NT treatments. A fertiliser Nitrophoska (12% N, 10% P, 10% K, and 1% S) was applied at the rate of 120 kg/ha for fodder maize and 200 kg/ha for oats. This was followed by a second dose of 65 kg urea/ha to all the treatments. Aitchison seed drill (Seedmatic 1112) was used for sowing and fertiliser application. At maturity sheep grazed both crops.

2.2. Soil sampling

Within each treatment, five cores (25 mm diameter) were taken randomly from each depth 0–100 and 100–200 mm depths and in each replicate, before spring cultivation and sowing. The cores were pooled, air-dried and sieved (<2 mm). Subsamples of the sieved soil were air-dried and used for pH, C and N analyses.

2.3. Soil analyses

Air-dried samples were used for the total C and N analyses. Total C in soils was analysed by a combustion method (Induction Furnace, Leco, St. Joseph, Mich.). Total soil N was determined by semi-micro Kjeldahl digestion followed by NH_4^+ -N measurements in the digests by an autoanalyser procedure.

Soil pH (1:2.5 water) was determined according to Blakemore et al. (1987).

SWC and temperature were measured for the 0–100 mm depth each time N_2O was measured. Soil samples were collected, weighed, oven-dried to constant mass at 105 °C, and reweighed. The final mass M_s , and the differences between fresh and dry masses M_w were used to calculate the gravimetric SWC:

$$\text{SWC} = \left(\frac{M_w}{M_s} \right) \times 100 \quad (1)$$

Soil temperature was measured with a handheld digital thermometer. A temperature probe was inserted into the soil next to each chamber and temperatures were recorded manually.

Soil characteristics of the site are shown in Table 1. Monthly rainfall data from 1996 to 1999 is shown in Fig. 1. Soil water (Fig. 2) and temperature at 100 mm depth were measured at each sampling time.

The cumulative rainfall during the study period was below the average rainfall received in most years. Rainfall distribution was particularly low during summer (December–February), with only half the normal rainfall for this season. The rainfall during the remainder of the study period was 10–20% lower than the 30-year mean monthly values. The mean monthly air temperatures for the summer months ranged between 15 and 20 °C, and coolest temperatures were recorded in the winter months of June–August (7–10 °C). The annual fluctuations in soil temperature at 100 mm depth followed similar trends to the air temperatures, and can be considered normal for this region.

Table 1
Soil characteristics of the experimental site (Massey University research site)^a

Characteristic	Depth (mm)	PP	CT	NT
Total C (g/kg)	0–100	27.50	21.70	24.30
	100–200	18.40	19.30	18.00
Total N (g/kg)	0–100	2.40	2.00	2.20
	100–200	1.70	1.80	1.70
pH (H ₂ O)	0–100	5.31	5.30	5.10
	100–200	5.32	5.26	5.10

PP: permanent pasture; CT: conventional tillage; NT: no-tillage.

^a Soil type: Ohakea silt loam (Gleyic luvisol).

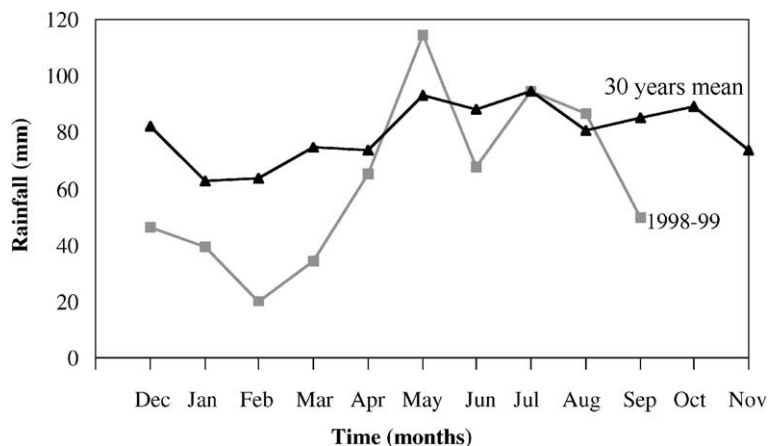


Fig. 1. Monthly rainfall from December 1998 to September 1999 and 30-year mean monthly rainfall.

2.4. Measurement of field N_2O emission

In order to assess the effect of tillage treatments on soil N_2O emissions, monthly (and occasionally fortnightly) measurements were taken during cropped as well as fallow periods. Fortnightly measurements were carried out when there was considerable weather change. The measurements at the Massey site started on 22 December 1998 and continued until 8 September

1999. During this period, a total of 10 measurements were completed. N_2O measurements in May were, inadvertently, not able to be taken.

Nitrous oxide emissions were measured following the chamber technique developed by Mosier and Hutchinson (1981). The chambers were modified PVC 'Sewer-hatch' (250 mm diameter and 300 mm high) attached to 200 mm deep and 250 mm diameter sections of PVC pipe. The PVC rim had an internal

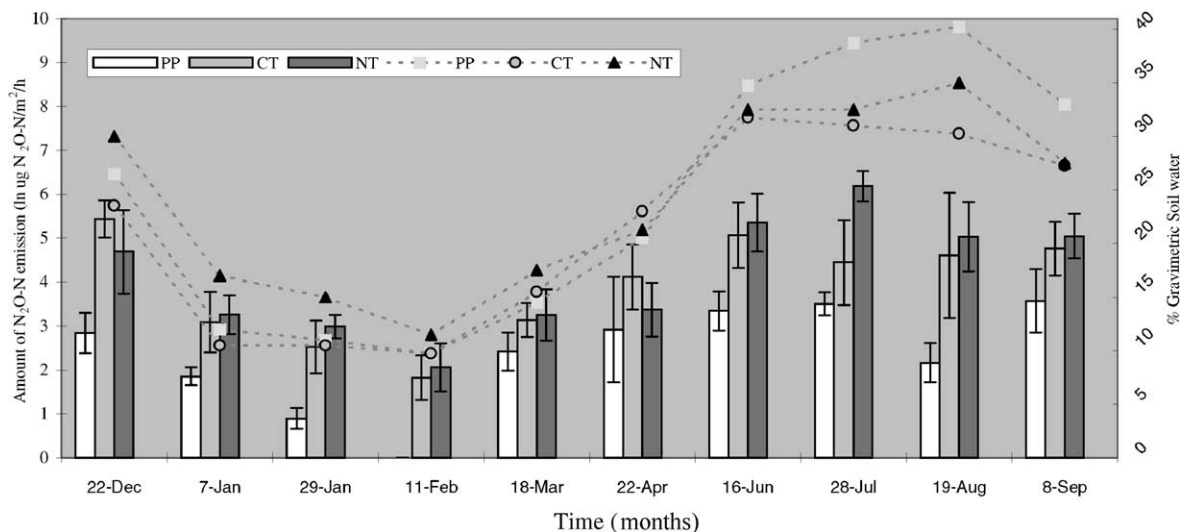


Fig. 2. Means of log-transformed (to the base e) nitrous oxide emissions ($\mu\text{g } N_2O\text{-N}/(\text{m}^2 \text{h})$), and gravimetric SWC. Vertical bars represent nitrous oxide flux and broken lines represent SWC. Vertical line bars show LSD values at $P < 0.05$.

half-turn locking system and a greased rubber 'O'-ring, to form an airtight chamber. At each measurement, 12 chambers, one chamber per plot (four replicates), were randomly inserted 100 mm into the soil 2 h before the measurements. Chamber heights were measured and each chamber volume calculated. The chambers were then sealed and headspace gas samples were collected using 60 ml polypropylene syringes fitted with three-way stopcocks. Because of concern over the permeability of polypropylene to N₂O, field samples were analysed as soon as possible after collection, usually within 24 h. Initially the chambers were left inserted in the soil for three hours with 1 h interval for air sampling in order to determine linear flux of N₂O production and to detect any leakage. In later samplings, it was reduced to only 2 h. Analyses were done using a gas chromatographic method (Mosier and Mack, 1980). Argon-methane carrier gas consisting of 10 ± 3% methane, <10 ppm oxygen and <5 ppm water in argon was used in the columns. Gas samples from syringes were introduced into a 5 cm³ gas sampling loop. Each analysis took approximately 6 min. Peaks were integrated with a Hewlett-Packard 3385A integrator. Ambient air at *t*₀, collected at the time of chamber installation, was used as a reference for calculating N₂O gas fluxes.

Two pilot experiments were conducted to assess (i) inherent spatial variations in N₂O fluxes due to soil heterogeneity (ii) the immediate effect of soil disturbance on N₂O fluxes during land preparation, after grazing fodder maize in March. N₂O fluxes were first measured for few days immediately after ploughing and subsequent measurements were taken, immediately after power harrowing. To determine if a possible relationship existed between N₂O fluxes and gravimetric SWC at 0–100 mm depth, water measurements were made on samples taken close to the installed chambers at the same time as gas emission samples were taken.

2.5. Statistical analysis

Soil analysis results were expressed on the basis of the oven-dry (105 °C) weight of the material. A general linear model (GLM) procedure was used for analysis of experimental data. An analysis of variance (ANOVA) using test of least significant difference (LSD) at 5% confidence level was used for comparisons of treatments. Correlation of SWC and nitrous oxide fluxes was analysed by a mixed model procedure.

3. Results and discussion

3.1. Spatial variability

The first pilot experiment showed there was a large inherent variation in N₂O fluxes with coefficient of variation values ranging between 110 and 127% across all treatments (Table 2), which reflected natural soil heterogeneity, and perhaps the measurement technique used rather than the real differences due to the tillage and cropping systems evaluated. Such variability in N₂O flux values is not unexpected, as Carran et al. (1995) have also found large field variability in the N₂O emissions because of soil heterogeneity and the episodic nature of N₂O emissions. A recent study by Kessavalou et al. (1998) showed that to accurately estimate annual net greenhouse gas fluxes for an agro-ecosystem, gas flux determinations must also be made at critical times near tillage and wetting events in addition to monitoring fluxes under normal field conditions.

The second pilot study (data not presented) showed significant differences in N₂O emissions between two sampling occasions. N₂O fluxes measured immediately after mouldboard ploughing but before power harrowing were 64 ± 14 µg N₂O-N/(m² h) but these

Table 2
Nitrous oxide emissions (µg N₂O-N/m² per day) from 12 replicate chambers showing spatial variation

Treatment	R1	R2	R3	R4	R5	R6	R7	R8	R9	R10	R11	R12	Mean	CV (%)
PP	456	134	144	696	43	62	166	67	790	67	24	36	223	120
CT	86	91	221	163	247	218	410	156	1010	554	1594	168	413	110
NT	341	2342	173	415	485	146	142	1313	300	391	96	91	518	127

PP: permanent pasture; CT: conventional tillage; NT: no-tillage.

dropped to $24 \pm 8 \mu\text{g N}_2\text{O-N}/(\text{m}^2 \text{h})$ when measured immediately after power harrowing. Comparison of two events showed those N_2O fluxes declined by over 65% within 1 h after power harrowing. A similar observation was recorded by Kessavalou et al. (1998) during summer period, when N_2O emissions declined by about 64% within 30 min following disking. These authors also noted that, although N_2O emissions returned to the pre-disking level 8 h after disking, they declined again and remained low for the rest of the study period. The authors related these phenomena to low water filled pore space (WFPS) conditions in the top 7.6 cm of soil. In the present study, N_2O emission rate returned to the pre-power harrowing levels at the next sampling period, 1 month after power harrowing, which most likely was associated with increased SWC.

There were marked spatial variations in N_2O emissions in all treatments throughout the study period (Fig. 2). Cumulative CVs of treatments ranged from 39 to 140%. This was not unexpected, as other researchers, such as Lemke et al. (1998) have reported up to 92% of the large-scale spatial variability, which the authors related to the differences in soil clay content in that experimental site. Others reported more than two times higher spatial variations in N_2O data (Mahmood et al., 1998; Teira-Esmatges et al., 1998). Clayton et al. (1994) noted that N_2O emissions arising from denitrification tended to be more spatially variable than those arising from the aerobic process of nitrification are. Incubations of soil with acetylene (data not reported) suggested that large N_2O fluxes at the Kairanga field site can also result in high denitrification potential.

The spatial variability in denitrification could be attributed to "hot-spots" associated with high local concentrations of organic matter, nitrate level and soil condition. As Christensen and Tiedje (1998) have observed, the formation of these hot-spots was governed by the creation of anaerobic conditions due to increased respiration, following the introduction of a source of decomposable organic matter to the soil microbes. Clemens et al. (1999) have also reported similar observations.

Randomly selected sampling positions within the field could further contribute to large variability in N_2O fluxes. On some sampling occasions N_2O fluxes ranged almost 10-fold with the lowest and highest

recorded within the same treatment. As the plots were at approximately 4° slope, the lower end of the field tended to have more biological activity since fluxes from this end were generally approximately two times greater than from the upper end of the field.

3.2. Tillage effect on field N_2O emissions

The lowest and the highest fluxes ranged widely from 1 to $7.41 \mu\text{g N}_2\text{O-N}/(\text{m}^2 \text{h})$, respectively. Although soils sometimes act as a sink (Mahmood et al., 1998; Teira-Esmatges et al., 1998) there was no evidence of negative N_2O fluxes in the chambers. Statistical analyses using raw data showed non-normality of error distribution and comparison of treatment means did not show real differences. The N_2O emission rates were found to be log-normally distributed. Log-normally distributed N_2O emission rates have also been reported by Ambus and Christensen (1994).

Further computation and interpretations of the log-transformed data highlighted differences in some samples whereas initial analysis indicated no such differences (Fig. 2). On most sampling dates, the NT and CT treatments had significantly higher N_2O fluxes than those from the PP treatment. The March and April measurements showed no such difference among the treatments, which appeared to be mainly associated with high variability among the replicates in each treatment. Differences in emission values again pointed to high variability in the N_2O emission data.

On an annualised basis, N_2O emissions measured from December 1998 to September 1999 from the PP ($1.66 \text{ kg N}_2\text{O-N}/\text{ha}$ per year or $19 \mu\text{g N}_2\text{O-N}/(\text{m}^2 \text{h})$) were significantly lower than the CT and NT fields averaging at 9.20 (or 105) and 12.0 (or 137) $\text{kg N}_2\text{O-N}/\text{ha}$ per year (or $\mu\text{g N}_2\text{O-N}/(\text{m}^2 \text{h})$), respectively (Table 3). However, there were no differences in N_2O emission rates between the CT and NT treatments. The data on seasonal N_2O emissions from PP, CT and NT is summarised in Table 4. The N_2O emissions during the fodder maize growth period (January–March) were generally low (see log-transformed data in Fig. 2). Starting from December when fluxes were high, with the PP, CT and NT treatments at 19, 246 and $151 \mu\text{g N}_2\text{O-N}/(\text{m}^2 \text{h})$, respectively, N_2O emissions declined considerably culminating in the lowest emissions in all treatments

Table 3
Effects of tillage systems on the means of N₂O emissions

Treatment	N ₂ O emissions/h ($\mu\text{g N}_2\text{O-N/m}^2$)	N ₂ O emissions per day ($\mu\text{g N}_2\text{O-N/m}^2$)	N ₂ O emissions per year (kg N ₂ O-N/ha)
PP	19 b	456 b	1.66 b
CT	105 a	2520 a	9.20 a
NT	137 a	3290 a	12.00 a

Values with the same letter in each column are not significantly different ($P < 0.05$). PP: permanent pasture; CT: conventional tillage; NT: no-tillage.

in February with the PP, CT and NT at 1, 7 and 9 $\mu\text{g N}_2\text{O-N}/(\text{m}^2 \text{h})$, respectively.

It is worth noting that despite the particularly low rainfall during summer (December–February), December N₂O emissions were several folds larger than those measured during January and February (Fig. 2) as the soil moisture in December was high. A similar phenomenon was recorded by Carran et al. (1995) who attributed it to the relatively large pool of both NH_4^+ and NO_3^- in the soil.

This sharp decline coincided, firstly, with dry soil water conditions, which had dropped by more than half compared to December SWC (Fig. 2). This corresponded with low rainfall in February, which was only 20.1 mm for the month. Secondly, lowest emissions also coincided with low rainfall and high temperatures during the 1998/1999 summer (Fig. 1), which are some of the key factors controlling denitrification. Nevertheless, N₂O fluxes from the CT and NT plots always remained higher than those from the PP plots.

From February onward, N₂O emission rates started to gradually increase. A few days after tillage in first the week of April, gas emissions from the CT were increased significantly compared to the NT and PP plots (data not reported). Although N₂O flux in the CT was 54 and 81% higher than that in the NT and PP, respectively a day earlier, it declined sharply after

seedbed preparation with power harrow. These results showed likely effects on N₂O efflux of recently cultivated soil as earlier suggested by Kessavalou et al. (1998).

Gas emissions following the tillage operation and during oats crop establishment in April showed relatively, although insignificantly, higher N₂O flux from the CT (55 and 62%) as compared to the NT and PP plots, respectively. These data suggest that cultivated soil allow higher or similar N₂O emissions to the NT soil during autumn season. Spring fluxes from CT were also found to be higher by MacKenzie et al. (1998). The N₂O emissions (Fig. 2) from June to September (during winter and spring period) generally reflected increased microbial activity irrespective of tillage or cropping system as reported by Aslam et al. (1999). The N₂O emissions during this period were consistently high reaching up to 684, 741 and 69 $\mu\text{g N}_2\text{O-N}/(\text{m}^2 \text{h})$ in the CT, NT and PP treatments, respectively (see log-transformed data in Fig. 2). The PP, however, continued to emit lower amounts of N₂O gas as compared to the NT and CT treatments even with the increased SWC. These findings support earlier suggestions by a number of researchers (Kaiser et al., 1998; Lemke et al., 1998; MacKenzie et al., 1998) who postulated that N₂O emissions were significantly affected by climate.

Table 4
Seasonal field N₂O emissions as affected by the PP, CT and NT treatments

Treatment	N ₂ O emissions (kg N ₂ O-N/ha)			
	Summer (December–February)	Autumn (March–May)	Winter (June–August)	LSD _{0.05}
PP	0.22 b	0.29 a	0.55 a	0.12
CT	1.61 b	1.53 b	2.91 a	0.17
NT	1.15 b	0.70 b	6.85 a	0.47

Values with the same letter in each row are not significantly different ($P < 0.05$). PP: permanent pasture; CT: conventional tillage; NT: no-tillage.

The New Zealand winter is marked by low soil temperatures and high rainfall. This resulted in increased SWC. Despite low temperatures, high soil water increased N₂O emissions significantly as a result of denitrification. Because N₂O is a product of both nitrification and denitrification processes, nitrification was presumably the major source of summer N₂O flux, as SWC during this period decreased and allowed predominantly aerobic soil conditions. Conversely, high N₂O fluxes during winter (Table 4) indicated that denitrification was the major source of N₂O emissions from these soils.

3.3. Effect of soil water content (SWC) on N₂O emissions

Tillage practices generally influenced water content in all treatments during the year. The NT plots had higher SWC compared to the CT and PP plots during summer and early autumn when rainfall events were not frequent. Lower water content in the PP could be due to high transpiration rate from the growing pasture. However, during the winter period soil water in the PP was generally higher compared to the NT and CT plots that could be associated with generally low evapotranspiration.

Generally, SWC appeared to govern N₂O emission rates. The lowest fluxes occurred when SWC was low (during summer period) and tended to increase with higher SWC during the winter period (Fig. 2). A number of regressions analyses to obtain the best-fit equation between the SWC and monthly N₂O fluxes were computed. To examine these relationships further transformations were computed to observe if any relationship existed between the values of N₂O fluxes and soil water (Fig. 3).

Overall, the log-transformed N₂O data showed better correlation with field soil water with the PP ($r = 0.73$), CT ($r = 0.75$) and NT ($r = 0.86$). This indicated that SWC and N₂O emissions had strong relationship in all treatments. Nevertheless, a large number of samples were required to fully explain the relationship between the soil water and N₂O emission rates. These relationships indicated that field SWC generally, although partially, determined the N₂O emissions.

Other factors, such as soil temperature also played a role in the amount of N₂O release. These data although not shown imply that in the Manawatu region of New

Zealand, the onset of rainy period is a major factor triggering N₂O emissions.

3.4. Effects of seasonal variations on N₂O emissions

Regardless of cropping and tillage management used, general seasonal patterns at this site are well defined, although 1 year data was not enough to statistically build time series. Low N₂O emission values were predominant during the dry season while wet season values were considerably higher (Fig. 2). This trend was similar across all tillage treatments.

Seasonal grouping of N₂O data showed significant differences in the N₂O emissions in all three treatments (Table 4). Because there was only one measurement conducted in the first week of September, spring season was not included in the comparison. Seasonal variation in N₂O emission from the PP was in the order of winter > autumn = summer. Similarly, although fluxes in the CT were higher in winter than in the autumn season, there were no differences between the summer and autumn data. The seasonal variations in N₂O emission from the NT treatment were also in the order of winter > autumn = summer.

The seasonal variations (Table 4) indicated that N₂O emissions were generally higher during winter as compared to summer and autumn. This could be triggered by wet weather, which was more frequent and intense generally starting from the end of autumn and continuing during winter, thus creating favourable conditions for denitrification.

The data also showed large seasonal as well as monthly variation in N₂O emissions. Soil water and temperature presumably caused some of this variability. For example, N₂O fluxes during the first sampling on 22 December were several orders higher than the next sampling on 7 January (Fig. 2). This suggested that high soil temperature coupled with the major rainfall event in December possibly gave bursts of N₂O fluxes. Kessavalou et al. (1998) have also reported up to five-fold increase in N₂O emissions following wetting. These authors had postulated that failure to include these short-lived episodic gas pulses in annual flux estimations might underestimate annual N₂O loss by up to 24%.

It is worth noting that when considering the 30-year rainfall data (Fig. 1) during measurement period, it showed a clear deviation in 1998–1999 rainfall. The

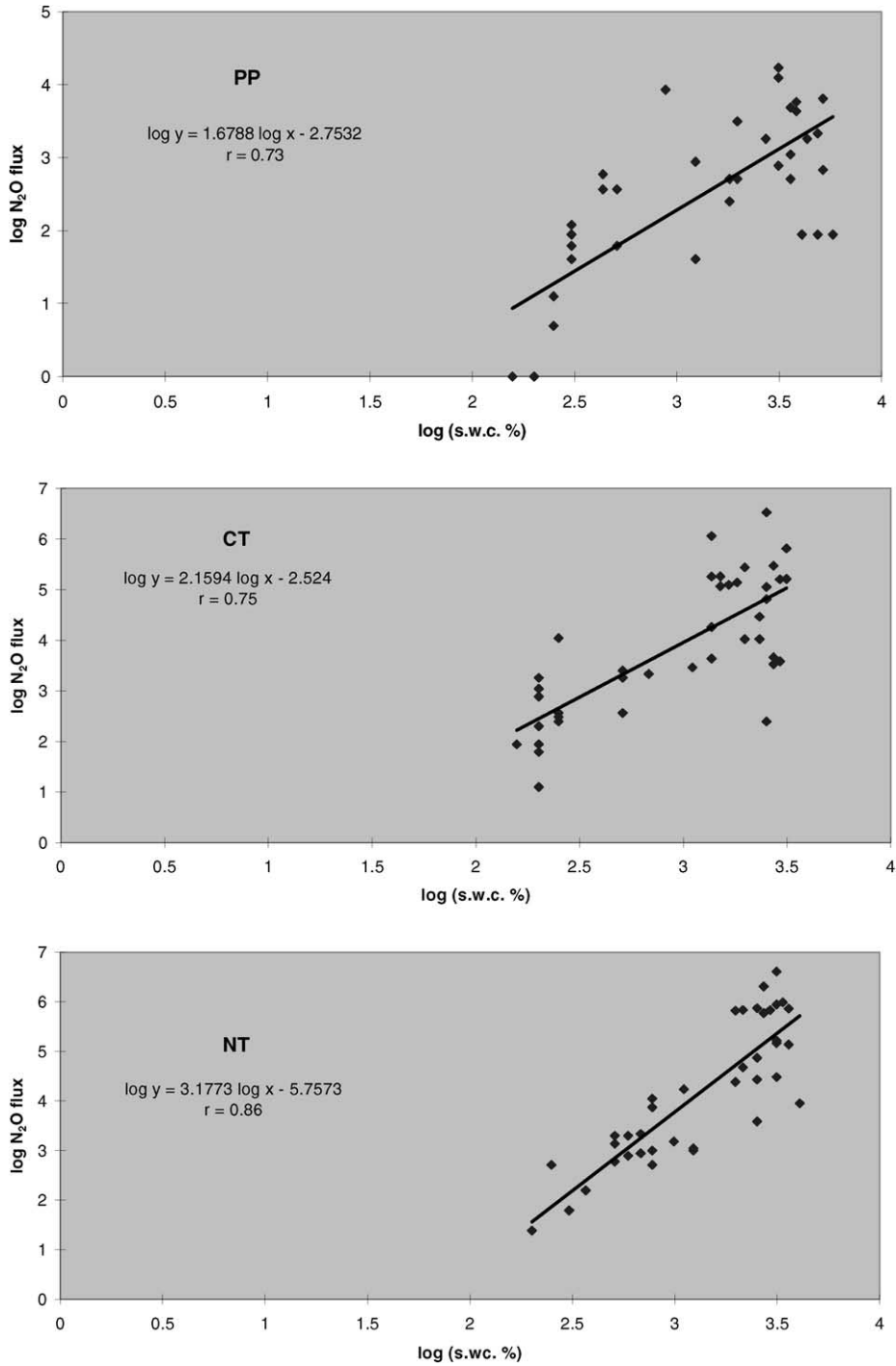


Fig. 3. Regression analysis of log-transformed (to the base e) data between SWC and N₂O emissions ($\mu\text{g N}_2\text{O-N}/(\text{m}^2 \text{h})$). PP: permanent pasture; CT: conventional tillage; and NT: no-tillage; SWC: soil water content.

cumulative rainfall during the study period was below the average rainfall received in 30 years. It was particularly low during summer (December–March) and early spring (September), with only half the normal rainfall for this season. The rainfall was 10–20% lower during winter (April and June), and ~20% higher in May than the 30-year mean monthly values. This may have had an effect on time of release and the flux magnitude of N₂O. Kaiser et al. (1998) found approximately 50% of the annual losses during winter period from October to February in the northern hemisphere. This highlights the importance of this period for the assessment of total N₂O losses from arable lands in the temperate climate of New Zealand.

4. Conclusions

There was large spatial variability in the N₂O data when measured in the field because of soil heterogeneity, and perhaps the measurement technique used rather than the real differences due to the tillage and cropping systems evaluated.

On an annualised basis, N₂O emissions measured from December 1998 to September 1999 from the PP (1.66 kg N₂O-N/ha per year) were significantly lower than the CT and NT fields averaging at 9.20 and 12.0 kg N₂O-N/ha per year, respectively. The N₂O emissions measured from this occasionally sheep grazed PP were one-half of those reported for intensively grazed by dairy stock (Choudhary et al., 2001).

Nitrous oxide fluxes were substantially higher in the NT and CT compared to the PP treatment. However, there were no marked differences between CT and NT treatments. This indicates that conversion of pastureland to cropping significantly increases N₂O emissions. The possible reason of rather low N₂O emissions from the PP could be the greater nitrogen utilisation by the growing grasses, whereas in the CT and NT fields the root distribution is not as dense as proliferation of grass roots. This suggested that perhaps considerably less nitrification occurred in grassland soils than in cropped soils, allowing less opportunity for denitrification in grassland soils.

The strong correlation between N₂O emissions and SWC in all treatments suggests that the high rainfall and wet winter and early spring together with soil properties, such as drainage characteristics are

important in the assessment of N₂O fluxes from these fields. Therefore, mitigation practices to reduce N₂O emissions should be directed towards the periods of high emissions. Regardless of cropping and tillage management, general seasonal patterns at this site were obvious, although 1 year data was not enough to build statistical time series. Further extended studies with a high frequency of sampling are, therefore, necessary to confirm and obtain a robust average yearly value for New Zealand pastures grazed by sheep, sheep-beef and dairy cattle, and CT and NT soils.

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